

# COMPUTATIONAL GEODYNAMICS FOR SEISMIC HAZARD ANALYSIS AND EARTHQUAKE PREDICTION

**Alik Ismail-Zadeh**

*Institute of Applied Geosciences, Karlsruhe Institute of Technology, Karlsruhe, Germany  
Institute of Earthquake Prediction Theory and Mathematical Geophysics, Russian Academy of Sciences, Moscow, Russian Federation  
Institut de Physique du Globe de Paris, Paris, France*

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## Summary

Recent advances in understanding Earth dynamics and in development of computational tools permit accurate numerical modeling and forecasting that are transforming the Earth sciences. These advances have a strong impact on studies of lithosphere dynamics, tectonic stress localization and changes, comprehensive seismic hazard assessment, and show significant potential to be applied to serve the sustainable development of society.

The vulnerability of human civilization to disasters caused by large earthquakes is growing due to the clustering of populations and proliferation of high-risk objects. The disasters have become a threat to a civilization's well-being: today a single earthquake may take up to several hundred thousand lives and cause significant material damage. The problem of earthquake prediction is a "grand challenge", although it may be solvable.

The ability to successfully predict large earthquakes would provide society with a change of reducing damage and loss of life due to ground motion, landslides, fire,

tsunamis and other associated hazards, and would assist in disaster preparedness protecting a society's population, economy, and environment. This chapter will introduce the reader to computational geodynamics and discuss its application to seismic hazard analysis. After a brief introduction into the subject, essential topics for comprehensive hazards assessment will be discussed. Special emphasis is placed on tectonic stress modeling, earthquake simulation, and seismic hazards assessment. The importance of quantitative seismic risk assessment and earthquake prediction studies are then discussed, and some final remarks are made in the conclusion.

## **1. Introduction to Computational Geodynamics**

*Geodynamics* deals with the dynamical, physical and chemical processes in the Earth's interior. Phenomena subject includes fundamental geological processes such as plate tectonics, mantle convection and their interaction. According to plate tectonic theory (e.g., Turcotte and Schubert, 2002), the plates are continually created and consumed. At oceanic ridges adjacent plates diverge in a process, which is known as seafloor spreading. As the lithosphere moves away from oceanic ridges, it cools, densifies, and thickens.

Once the lithosphere becomes sufficiently dense compared to the underlying mantle rocks, it bends, founders, and begins sinking into the hot mantle due to gravitational instability. Because the lithosphere behaves elastically at short time scales (seconds to years), it can transmit stresses. The downward buoyancy forces (generated due to the excess density of the rocks of the descending lithosphere) promote the sinking of the lithosphere, but elastic, viscous and frictional forces resist the descent.

The combination of these forces produces shear stresses high enough to cause earthquakes. Other processes contributing to stress generation in the descending lithosphere and its release in earthquakes can be plastic instability at high temperature, faulting due to metamorphic phase transitions, and dehydration-induced embrittlement. Oceanic trenches are the sites of the world largest earthquakes. The earthquakes at oceanic trench zones can occur along the descending lithosphere to the depths of about 660 km depending on the thermal state in the mantle.

Modern geodynamics was born in the late 1960's with the general acceptance of the plate tectonics paradigm. At the beginning, simple analytical models were developed to explain plate tectonics and its associated geological structures (see, e.g., Turcotte and Schubert, 2002, for the analytical models). These models were highly successful explaining many of the first order features of behavior of the Earth. The necessity to go beyond these basic models in order to make them more realistic and to understand better the Earth shifted the emphasis to numerical simulations.

The development of computers heralded a new era in the approach to the solution of geodynamical problems. Programming languages, operating systems, management of large quantities of data, correctness of numerical codes and many other considerations relevant to the efficient and accurate solution of the problems became subjects of the new discipline of computer science, on which scientific computing now depends heavily. Nonetheless, mathematics continues to play a major role in scientific

computing: it provides information about the suitability of a model and the theoretical foundation for the numerical methods. The numerical models of geodynamical problems have grown increasingly complex and capable over time with improvements in computational power and numerical algorithms. This has resulted in the development of a new branch of geoscience called *computational geodynamics*.

Characteristic of this new intellectual landscape is the need for strong interaction across traditional disciplinary boundaries: geodynamics, mathematics, and computer science. Mathematics provides the means to establish the credibility of numerical methods and algorithms, such as error analysis, exact solutions, uniqueness and stability analysis.

Computer science provides the tools, ranging from networking and visualization tools to algorithms matching modern computer architectures. The power of computational geodynamics is so huge that it can deal not only with the quantitative assessment of forces and processes in the Earth's interior, but also to contribute to the understanding and assessment of geohazards (e.g., seismic hazards) and assist in earthquake prediction.

### **1.1. Mathematical Approach to Geodynamic Problems**

Many geodynamic problems can be described by mathematical models, i.e. by a set of partial differential equations and boundary and/or initial conditions defined in a specific domain. Models in computational geodynamics predict quantitatively what will happen when the crust, lithosphere, and the mantle deform, often with the inclusion of complications due to factors such as simultaneous heat transport, phase changes in the Earth's interior, complex rheology, melting and melt migration, chemical reactions, solid body motion, etc.

Mathematical models of geodynamic processes can be solved analytically or numerically. Analytical solutions are those that a researcher can obtain by solving mathematical models using a pencil, a piece of paper, and thought. Simple mathematical models allow analytical solutions, which have been (and still are) of great importance because of their power: the solutions are precise and can be presented by exact formulas.

However, the usefulness of exactly solvable models is limited as many mathematical models of geodynamics are too complicated to be solved analytically. Numerical solutions are those that researchers can obtain by solving models using numerical methods and computers. Numerical models allow the solution of complex problems of geodynamic processes, although the solutions are not exact. In some geodynamic applications an analytical solution to part of the complex problem can be implemented into the numerical model to make the model much more effective.

### **1.2. Computational Approach to Geodynamic Problems**

Only a few of the ordinary and partial differential equations describing geodynamical models can be solved exactly, and hence the equations are transformed into discrete equations to be solved numerically. The widespread access to high-performance

computers has resulted in an over-reliance on numerical answers in spite of other possibilities, and has created a corresponding false sense of security with respect to serious numerical problems or errors. Nevertheless, it is now possible without too much trouble to find solutions to most equations that are routinely encountered.

According to Ismail-Zadeh and Tackley (2010), the rationale of the numerical modeling can be described as follows (see Figure 1). The initial stage of numerical modeling is to describe geodynamically complex reality by a simplification of the reality; namely, to introduce the model concept of the geodynamic problem, forces acting on the system (crust/lithosphere/mantle), physical parameters to be used in the modeling, etc. A physical model is then developed to which the physical laws can be applied. The next step in the numerical modeling is to describe the physical model by means of mathematical equations.

The comparison with observations allows the model to be tested (validated). If a mathematical model is found to be inadequate - the assumed process is not the correct one, or some significant factors have been missed - it must be refined. The mathematical model should be properly determined, at least after the numerical values of some still unknown parameters have been determined (that is, the model is tuned). Once the mathematical model is developed, proper numerical tools and methods have to be chosen, and relevant numerical codes (software) should be developed (or otherwise obtained).

The mathematical model should be transformed into the computational model containing discrete equations to be solved by using computers. An important element of numerical modeling is verification of the model, namely, the assessment of the accuracy of the solution to the computational model by comparison with known solutions (analytical or numerical). Once the computational model is verified, the numerical model can be used to simulate reality, and the numerical results can be tested against observations. If there is a good agreement between the numerical results and observed (field or experimental) data, the model results can be considered as valid model predictions.

An essential component of computational geodynamics is a *numerical method*. Although many of the key ideas for numerical solution methods were established several centuries ago, they were of little use before computers appeared. Interest in numerical methods increased dramatically with the development of the computer power. Numerical methods provide possibilities to obtain accurate solutions to geodynamic problems. However, the numerical results are always approximate.

There are reasons for differences between computed results and observations. Errors arise from each part of the process used to produce numerical solution: the physical model is too simplified compared to geodynamic reality; the equations (mathematical model) may contain approximations or idealizations; approximations are made in the discretization process; in solving the discrete equations, iterative methods are used and insufficient iterations are taken. Additionally, uncertainty in physical parameters can lead to differences between computed results and observations.

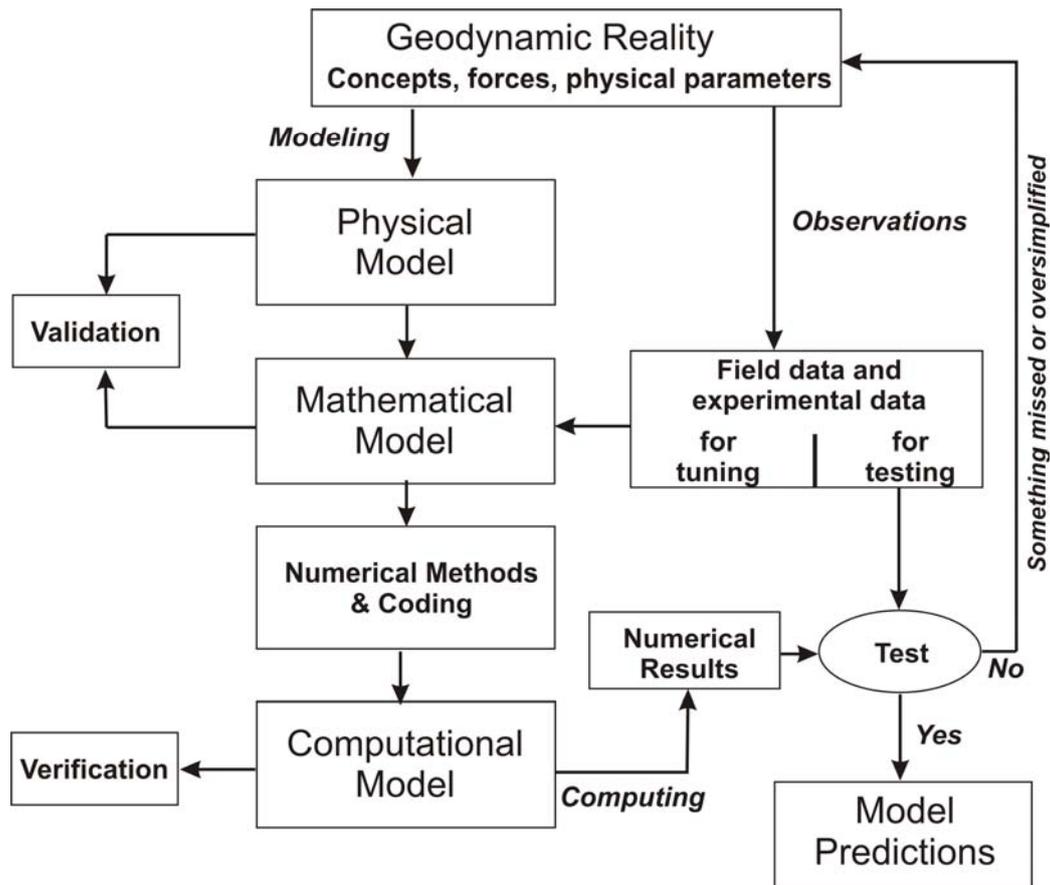


Figure 1. Schematic representation of the essential parts of numerical modeling (after Ismail-Zadeh and Tackley, 2010).

Numerical simulations in geodynamics enable one to analyze and to predict the dynamics of the Earth's interior. The basic elements of the numerical modeling are as follows: a mathematical model describing geodynamic processes; a discretization method to convert the mathematical equations into discrete equations to be solved numerically; numerical method(s) to solve the discretized equations; computer code(s) (i.e. software) to be developed or to be used, if already developed, which solve numerically the discrete equations; computer hardware, which performs the calculations; results of numerical modeling to be visualized, analyzed and interpreted by geoscientist(s).

## 2. Introduction to Seismic Hazards and Risk Analysis

*Seismic hazard* can be defined as a potentially damaging earthquake, which may cause the loss of life or injury, property damage, social and economic disruption, or environmental degradation. Seismic hazard is typically interrelated with past seismicity and geological and geophysical parameters (e.g. peak ground acceleration, seismic intensity, seismic wave propagation and attenuation, site effect). *Seismic risk* is a measure that combines the likelihoods and the consequences, over a given time, of a set of earthquake scenarios.

Natural hazard and risk analysis should be considered from a holistic point of view (from the whole to details). According to Ismail-Zadeh (2010), a holistic comprehensive quantitative assessment of seismic hazard should be based on multidisciplinary research in (i) geodynamics and geodesy (to identify zones of tectonic stress and strain localization), (ii) present and historical seismicity (to localize areas prone to strong events), (iii) nonlinear dynamics of the lithosphere (to analyze statistical properties of the earthquake sequences, their clustering and critical transitions), (iv) soil property (to analyze liquefaction and seismic shaking), and (v) standard hazards assessment (to determine peak ground acceleration, response spectra amplitude, and seismic intensity). This approach to seismic hazard should be accompanied by a holistic approach to earthquake prediction (e.g., Keilis-Borok and Soloviev, 2003) and by a holistic approach to seismic risk, when a convolution of hazard, vulnerability and exposure (as functions of space and time) should be viewed also from socio-psychological (e.g., resilience of community to extreme seismic events) and legislation (e.g., role of law in risk reduction) points of view (Beer and Ismail-Zadeh, 2003).

The *vulnerability* of human civilization to natural disasters is growing due to the proliferation of high-risk objects, clustering of populations, and destabilization of large cities. Today a single earthquake may take up to several hundred thousand lives, and cause material damage up to several hundred billion EURs (see Munich Re., 2012) with a possible chain reaction expanding to a world-wide financial crisis and economic depression.

A large earthquake in (or close to) Tokyo might result in a world financial crisis, because many Japanese companies, which invested considerable funds in foreign enterprises, will withdraw these funds to rebuild or to restore the city infrastructure after the disaster. A large earthquake can trigger an ecological catastrophe (e.g. Chernobyl-type calamities or recent Fukushima incident) if it occurs in close vicinity to a nuclear power plant built in an earthquake-prone area.

Extreme seismic events (e.g., the 1755 Lisbon, the 2004 Aceh-Sumatra, and the 2011 Great East Japan earthquakes) are a manifestation of the complex behavior of the lithosphere structured as a hierarchical system of blocks of different sizes. Driven by mantle convection these lithospheric blocks are involved in relative movement, resulting in stress localization and earthquakes.

Despite the lithosphere behaving as a large non-linear system, featuring instability and deterministic chaos, some integral empirical regularities emerge, indicating a wide range of similarity, collective behavior, and the possibility for earthquake prediction (e.g., Keilis-Borok et al. 2001). These great earthquakes, when they occur, are surprising, and society is poorly prepared to deal with them.

Protecting human life and property against earthquake disasters requires an uninterrupted chain of research and civil protection tasks (Ismail-Zadeh, 2010): from (i) an understanding of the physics of earthquakes, their analysis and monitoring, through (ii) interpretation, modeling, seismic hazard assessment, and earthquake prediction, to (iii) delivery of the scientific forecasts to local authorities, public awareness, preparedness, and preventive disaster management.

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Ismail-Zadeh A., Matenco L., Radulian M., Cloetingh S., Panza G. (2012a). Geodynamic and intermediate-depth seismicity in Vrancea (the south-eastern Carpathians): Current State-of-the-Art, *Tectonophysics*, 530-531, 50–79. [This article reviews geology, geodynamics, seismic, gravity and geoelectric studies, and discusses qualitative and quantitative geodynamic models, models of earthquakes and seismic hazards, and perspectives in future research.]

Ismail-Zadeh A.T., Le Mouél J.L., Soloviev A., Tapponnier P., Vorobieva I. (2007a). Numerical modeling of crustal block-and-fault dynamics, earthquakes and slip rates in the Tibet-Himalayan region. *Earth and Planetary Science Letters*, 258, 465–485. [This paper presents the results of the earthquake simulator (BAFD model) applied to the Tibet-Himalayan region.]

Ivan I.A., Enescu B.D., Pantea A. (1998). Input for seismic hazard assessment using Vrancea source region. *Romanian Journal of Physics*, 43, 619–636. [This paper discusses the input for seismic hazard assessment and presents models of seismic hazards due to intermediate-term earthquakes.]

Jiricek R. (1979). Tectonic development of the Carpathian arc in the Oligocene and Neogene. In: Mahel M. (ed.), *Tectonic Profiles Through the Western Carpathians*. Bratislava, Geologicky ustav Dionyz Stura, pp. 203–212. [This paper presents geological and tectonic evolutions of the region since 40 million years ago.]

Jordan T.H., Chen Y.-T., Gasparini P., Madariaga R., Main I., Marzocchi W., Papadopoulos G., Sobolev G., Yamaoka K., and Zschau J. (2011) Operational earthquake forecasting: State of knowledge and guidelines for utilization, *Annals of Geophysics*, 54 (4), doi: 10.4401/ag-5350. [This paper reviews the state of the art in earthquake prediction and forecasting with a special emphasis on the 2009 L'Aquila earthquake.]

Kantorovich L., Keilis-Borok V.I., Molchan G. (1973). Seismic risk and principles of seismic zoning. In: Keilis-Borok V.I. (ed.), *Computational and Statistical Methods for Interpretation of Seismic data*. Moscow, Nauka, pp. 3-20 (in Russian). [This paper introduces a mathematical background for seismic risk assessment and presents the basic principles of seismic zoning.]

Keilis-Borok V.I. (1990). The lithosphere of the Earth as a nonlinear system with implications for earthquake prediction. *Reviews of Geophysics*, 28, 19–34. [In this paper the lithosphere is viewed as a hierarchy of volumes from tectonic plates to grains of rock, and the relative movement of the volumes is realized to a large extent through earthquakes. The movement is controlled by a wide variety of independent processes, concentrated in the thin boundary zones between the volumes. These processes transform the lithosphere into a large nonlinear system, featuring instability and deterministic chaos. The possibility of intermediate-term earthquake prediction is discussed.]

Keilis-Borok V.I., Kossobokov V.G. (1990). Premonitory activation of earthquake flow: algorithm M8. *Physics of the Earth and Planetary Interiors*, 61, 73–83. [Many of large earthquakes occurred in different regions of the world are preceded by specific low-magnitude seismic activation, which was depicted by the algorithm M8. The algorithm presented in the paper is designed to alert the times of increased probability of large earthquakes.]

Keilis-Borok V.I., Soloviev A.A. (2003). *Nonlinear Dynamics of the Lithosphere and Earthquake Prediction*, 337 p. Heidelberg, Springer. [The book provides the reader with a holistic approach to earthquake modeling and prediction.]

Keilis-Borok V.I., Knopoff L., Rotwain I.M., Allen C.R. (1988). Intermediate-term prediction of occurrence times of strong earthquakes. *Nature*, 335, 690–694. [Pattern recognition procedures for infrequent events are adapted to the problem of identifying patterns of clustering of small- and intermediate-scale seismicity before large earthquakes. Identification procedures derived from analysis of large California and Nevada earthquakes yield a high success rate when applied to other parts of the world.]

Keilis-Borok V.I., Ismail-Zadeh A.T., Kossobokov V.G., Shebalin P.N. (2001). Non-linear dynamics of the lithosphere and intermediate-term earthquake prediction. *Tectonophysics*, 338(3-4), 247–259. [The paper reviews the development in modeling of the lithosphere dynamics and earthquake prediction and discusses new possibilities in improving the accuracy of predictions.]

Kelleher J., Sykes L., Oliver J. (1973). Possible criteria for predicting earthquake locations and their application to major plate boundaries of Pacific and Caribbean. *Journal of Geophysical Research*, 78, 2547–2585. [This study presents several possible criteria for forecasting the locations of large shallow earthquakes of the near future along major plate boundaries and for assigning a crudely determined rating to those forecasts.]

King G.C.P., Stein R.S., Lin J. (1994). Static stress changes and the triggering of earthquakes. *Bulletin of the Seismological Society of America*, 84, 935–953. [This paper examines the general problem of how one earthquake might trigger another and how changes in Coulomb stress conditions associated with one or more earthquakes may trigger subsequent events.]

Knopoff L. (1999). Earthquake prediction is difficult but not impossible. *Nature debates*. <http://www.nature.com/nature/debates/earthquake>. Accessed 25 March 2012. [The paper discusses the earthquake prediction possibilities and emphasize important additional constraints: a utilitarian constraint demands that the lower magnitude bound be appropriate to societal needs.]

Kossobokov V. (2006). Quantitative earthquake prediction on global and regional scales. In: Ismail-Zadeh A.T. (ed.), *Recent Geodynamics, Georisk and Sustainable Development in the Black Sea to Caspian Sea Region*. American Institute of Physics Conference Proceedings, vol. 825, Melville, New York, pp. 32–50. [This chapter discusses earthquake prediction experiments conducted by the author on global and regional scales.]

Kossobokov V.G., Keilis-Borok V.I., Smith S.W. (1990). Localization of intermediate-term earthquake prediction. *Journal of Geophysical Research*, 95(B12), 19763–19772. [This paper presents a simple algorithm to refine the locality in which the strong earthquake may be expected to occur.]

Lighthill J., ed. (1996). *A Critical Review of VAN: Earthquake Prediction from Seismic Electric Signals*, 376 pp. Singapore, World Scientific Publication. [This book contains the chapters, which review the VAN method of earthquake prediction.]

Lin J., Stein, R.S. (2004). Stress triggering in thrust and subduction earthquakes, and stress interaction between the southern San Andreas and nearby thrust and strike-slip faults, *Journal of Geophysical Research*, 109, B02303, doi:10.1029/2003JB002607. [This paper argues that key features of thrust earthquake triggering, inhibition, and clustering can be explained by Coulomb stress changes, and this is illustrated by a suite of representative models and by detailed examples.]

Linzer H.G. (1996). Kinematics of retreating subduction along the Carpathian arc, Romania. *Geology*, 24, 167–170. [The paper presents the reconstruction of retreating subduction along the Carpathians using kinematic axes and resultant vectors of displacement along the Carpathian arc and the Apuseni Mountains.]

Lorinczi P., Houseman G.A. (2009). Lithospheric gravitational instability beneath the southeast Carpathians. *Tectonophysics*, 474, 322–336. [In this study, the authors explore a dynamical model for the Vrancea region based on the idea of viscous flow of the lithospheric mantle permitting the development of local continental mantle downwelling beneath Vrancea, due to the gravitational instability.]

Lungu D., Cornea T., Nedelcu C. (1999). Hazard assessment and site dependent response for Vrancea earthquakes. In: Wenzel F., Lungu D., Novak O. (eds.), *Vrancea Earthquakes: Tectonics, Hazard and Risk Mitigation*. Dordrecht, Kluwer Academic Publishers, pp. 251–267. [The paper discusses seismic assessment and site response for intermediate-depth earthquakes in Vrancea.]

Lyakhovskiy V., Ben-Zion Y., and Agnon, A. (2001). Earthquake cycle, fault zones, and seismicity patterns in a rheologically layered lithosphere. *Journal of Geophysical Research*, 106, 4103–4120. [This paper presents the coupled evolution of earthquakes and faults in a model consisting of a seismogenic upper crust governed by damage rheology over a viscoelastic substrate.]

Mandrescu N., Radulian M. (1999). Macroseismic field of the Romanian intermediate-depth earthquakes. In: Wenzel F., Lungu D., Novak O. (eds.), *Vrancea Earthquakes: Tectonics, Hazard and Risk Mitigation*. Dordrecht, Kluwer Academic Publishers, pp. 163–174. [The paper discusses seismicity and macroseismic field studies related to the Vrancea earthquakes.]

Manea V.C., Manea M. (2009). Thermally induced stressed beneath the Vrancea area. In: Besutiu L. (ed.), *Integrated Research on the Intermediate-Depth Earthquake Genesis Within Vrancea Zone*. Bucharest, Vergiliu, pp. 172–183. [The paper presents a thermally induced stress model based on the temperatures derived from seismic velocity anomalies.]

Mantyniemi P., Marza V.I., Kijko A., Retief, P. (2003). A new probabilistic seismic hazard analysis for the Vrancea (Romania) seismogenic zone. *Natural Hazard*, 29, 371–385. [In this paper, the authors apply a probabilistic methodology to map specific seismic hazard induced by the Vrancea Intermediate-depth earthquakes.]

Martin M., Wenzel F., the CALIXTO working group (2006). High-resolution teleseismic body wave tomography beneath SE-Romania – II. Imaging of a slab detachment scenario. *Geophysical Journal International*, 164, 579–595. [In this paper, the authors present the relative *P*-wave velocity distribution of the lithosphere/asthenosphere system in the region. Smearing from strong crustal velocity anomalies into the upper mantle is successfully suppressed by traveltimes corrections with an *a priori* 3-D regional crustal velocity model.]

McKenzie D.P. (1972). Active tectonics of the Mediterranean region. *Geophysical Journal of the Royal Astronomical Society*, 30, 109–185. [This article examines fault plane solutions for earthquakes between the Mid-Atlantic ridge and Eastern Iran and shows that the deformation at present occurring is the result of small continental plates moving away from Eastern Turkey and Western Iran.]

Molchan G., Romashkova L. (2011). Gambling score in earthquake prediction analysis. *Geophysical Journal International*, 184, 1445–1454. [The paper expands parameterization of the gambling score and uses the M8 algorithm for earthquake prediction to illustrate difficulties of the approach in the analysis of the prediction significance.]

Moldoveanu C.L., Panza G.F. (1999). Modeling for microzonation purposes of the seismic ground motion in Bucharest, due to the Vrancea earthquake of May 30, 1990. In: Wenzel, F., Lungu, D., Novak, O. (Eds.), *Vrancea Earthquakes: Tectonics, Hazard, and Risk Mitigation*. Dordrecht, Kluwer Academic Publishers, pp. 85–97. [This paper presents a model of seismic ground motion due to the strong 1990 Vrancea earthquake.]

Moldoveanu C.L., Panza G.F. (2001). Vrancea source influence on local seismic response in Bucharest. *Pure and Applied Geophysics*, 158, 2407–2429. [The mapping of the seismic ground motion in Bucharest due to the strong Vrancea earthquakes is carried out using a complex hybrid waveform modeling method that allows easy parametric tests.]

Morat P., Le Mouél J.-L. (1987). Variation of the electrical resistivity of large rock samples with stress, *Geophysics*, 52, 1424–1430. [This paper presents a study related to variations of mechanical stresses inside the rock and variations of its electrical properties at different scales (from a centimeter to a kilometer) to see if some scaling laws could be experimentally established.]

Munich Re (2012). Review of natural catastrophes in 2011: Earthquakes result in record loss year; [http://www.munichre.com/en/media\\_relations/press\\_releases/2012/2012\\_01\\_04\\_press\\_release.aspx](http://www.munichre.com/en/media_relations/press_releases/2012/2012_01_04_press_release.aspx). Accessed 15 March 2012. [This article reviews the natural catastrophes occurred in 2011 and presents statistics of natural disasters.]

Naimark B.N., Ismail-Zadeh A.T., Jacoby W.R. (1998). Numerical approach to problems of gravitational instability of geostructures with advected material boundaries. *Geophysical Journal International*, 134, 473–483. [The paper presents a numerical approach for solving 2-D mantle flow problems, where the chemical composition changes abruptly across intermediate boundaries. The method combines a

Galerkin-spline technique with a method of integration over regions bounded by advected interfaces to represent discontinuous variations of material parameters.]

Ogata Y. (1988). Statistical models for earthquake occurrences and residual analysis for point processes. *Journal of the American Statistical Association*, 83, 9–27. [This article discusses several classes of stochastic models for the origin times and magnitudes of earthquakes.]

Oncescu M.C. (1984). Deep structure of the Vrancea region, Romania, inferred from simultaneous inversion for hypocenters and 3-D velocity structure. *Annals of Geophysics*, 2, 23–28. [This paper discusses the deep structure of the Vrancea region and presents a model of earthquake occurrence in the Vrancea region.]

Oncescu M.C., Bonjer K.P. (1997). A note on the depth recurrence and strain release of large Vrancea earthquakes. *Tectonophysics*, 272, 291–302. [The large 1940 Vrancea event is investigated in this paper in terms of the depth of rupture initiation and extent. The ruptured zones of the 1977, 1986 and 1990 Vrancea events are estimated from the distribution of aftershocks located with the joint hypocenter determination method.]

Oncescu M.C., Bonjer K.P., Rizescu M. (1999). Weak and strong ground motion of intermediate depth earthquakes from the Vrancea region. In: Wenzel, F., Lungu, D., Novak, O. (Eds.), *Vrancea Earthquakes: Tectonics, Hazard, and Risk Mitigation*. Dordrecht, Kluwer Academic Publishers, pp. 43–47. [This article discusses the various types of ground motions due to Vrancea deep seismicity.]

Oxford University Press (2007). *The Oxford Dictionary of Synonyms and Antonyms*, 2<sup>nd</sup> Edition: Oxford University Press, Oxford, 528 p.

Panza G.F., Irikura K., Kouteva M., Peresan A., Wang Z., Saragoni R., eds. (2011). Advanced Seismic Hazard Assessment, *Pure and Applied Geophysics*, 168, DOI 10.1007/s00024-010-0179-9. [This volume presents multifaceted information on the modern tools for seismic hazard assessment, and to make clear the significant difference between hazard and risk, and hazard mitigation and risk reduction.]

Pollitz F.F. (2009). A viscoelastic earthquake simulator with application to the San Francisco bay region. *Bulletin of the Seismological Society of America*, 99, 1760–1785. [This paper presents an earthquake simulator based on elastic dislocation theory accounting for the effects of interseismic tectonic loading, static stress steps at the time of earthquakes, and post-earthquake stress readjustment through viscoelastic relaxation of the lower crust and mantle.]

Pollitz F.F., Sacks, I.S. (2002). Stress triggering of the 1999 Hector Mine earthquake by transient deformation following the 1992 Landers earthquake. *Bulletin of the Seismological Society of America*, 92, 1487–1496. [By employing a viscoelastic model calibrated by geodetic data collected during the time period between the Landers and Hector Mine events, the authors calculate that postseismic relaxation produced a transient increase in Coulomb failure stress of about 0.7 bars on the impending Hector Mine rupture surface.]

Radulian M., Vaccari F., Mandrescu N., Panza G.F., Moldoveanu C.L. (2000). Seismic hazard of Romania: deterministic approach. *Pure and Applied Geophysics*, 157, 221–247. [The seismic hazard is estimated in this paper in terms of peak-ground motion values-displacement, velocity, and design ground acceleration-computing complete synthetic seismograms.]

Reid H.F. (1911). The elastic-rebound theory of earthquakes. *University of California Publications in Geological Sciences*, 413-444. [This paper presents a theory of earthquake occurrence based on elastic rebound.]

Rundle J.B., Tiampo K.F., Klein W., Martins J.S.S. (2002). Self-organization in leaky threshold systems: The influence of near-mean field dynamics and its implications for earthquakes, neurobiology, and forecasting. *Proceedings of the National Academy of Sciences of the United States of America*, 99 (Suppl. 1), 2514–2521. [This paper discusses the physics of self-organization in earthquake threshold systems at distinct scales.]

Rundle P.B., Rundle J.B., Tiampo K.F., Donnellan A., Turcotte D.L. (2006). Virtual California: Fault model, frictional parameters, applications, *Pure and Applied Geophysics*, 163, 1819–1846. [The Virtual California is a topologically realistic simulation of the interacting earthquake faults in California. Inputs to the model arise from field data, and typically include realistic fault system topologies, realistic long-term slip rates, and realistic frictional parameters.]

Schwartz D.P., Coppersmith K.J. (1984). Fault behavior and characteristic earthquakes - examples from the Wasatch and San Andreas fault zones, *Journal of Geophysical Research*, 89, 5681–5698. [This paper studies earthquake recurrence relationships on the Wasatch and San Andreas faults based on historical seismicity data and geologic data and shows that a linear extrapolation of the cumulative recurrence curve from the smaller magnitudes leads to gross underestimates of the frequency of occurrence of the large or characteristic earthquakes.]

Shebalin P., Kellis-Borok V., Gabrielov A., Zaliapin I., Turcotte D. (2006). Short-term earthquake prediction by reverse analysis of lithosphere dynamics. *Tectonophysics*, 413, 63-75. [This paper describes a methodology for short-term prediction named RTP (Reverse Tracing of Precursors).]

Sobolev G. (2001). The examples of earthquake preparation in Kamchatka and Japan. *Tectonophysics*, 338, 269–279. [The paper presents the algorithm for earthquake prediction and the results of testing of the algorithm in Kamchatka and Japan.]

Sokolov V.Y, Bonjer K.-P., Wenzel F. (2004). Accounting for site effect in probabilistic assessment of seismic hazard for Romania and Bucharest: A case of deep seismicity in Vrancea. *Soil Dynamics and Earthquake Engineering*, 24, 929–947. [The paper presents the evaluations of site-dependent seismic hazard in Romania and in the capital city of Bucharest caused by the Vrancea intermediate-depth seismicity.]

Sokolov V.Y, Bonjer K-P., Oncescu M., Rizescu M. (2005). Hard rock spectral models for intermediate-depth Vrancea (Romania) earthquakes. *Bulletin of the Seismological Society of America*, 95, 1749–1765. [The frequency-dependent amplification for rock sites and apparent source spectra are studied using an earthquake ground-motion database collected in Romania.]

Soloviev A.A., Ismail-Zadeh A.T. (2003). Models of dynamics of block-and-fault systems. In: Keilis-Borok V.I., Soloviev A.A. (eds.), *Nonlinear Dynamics of the Lithosphere and Earthquake Prediction*. Springer, Heidelberg, pp. 69–138. [This chapter describes the BAFD model of seismicity and presents applications of the model to several earthquake-prone regions.]

Sperner B., Lorenz F., Bonjer K., Hettel S., Müller B., Wenzel F. (2001). Slab break-off – abrupt cut or gradual detachment? New insights from the Vrancea region (SE Carpathians, Romania). *Terra Nova*, 13, 172–179. [Slab break-off is studied in the paper in correlation with the intermediate-depth seismic activity in the Vrancea region and with the low seismicity in the lowermost crust and uppermost mantle.]

Sykes L.R., Shaw B.E., Scholz C.H. (1999). Rethinking earthquake prediction. *Pure and Applied Geophysics*, 155, 207–232. [The paper summarizes what was possible by end of 20<sup>th</sup> century in predicting earthquakes, what might be accomplished and hence might be possible in the next few decades, and what types of predictions appear to be inherently impossible based on our understanding of earthquakes as complex phenomena.]

Turcotte D.L., Schubert G. (2002). *Geodynamics*, 2nd edn, 456 pp., Cambridge, Cambridge University Press. [This is a comprehensive textbook on geodynamics.]

Turcotte D.L., Holliday J.R., Rundle J.B. (2007). BASS, an alternative to ETAS. *Geophysical Research Letters*, 34, L12303, doi:10.1029/2007GL029696. [In this paper, the authors introduce the branching aftershock sequence (BASS) model, which utilizes Båth's law.]

Utsu T., Seki A. (1954). A relation between the area of aftershock region and the energy of main shock, *Journal of Seismological Society of Japan*, 7, 233–240. [This work presents the relationship between the magnitude of earthquake and the area of rupture.]

Uyeda S., Hayakawa M., Nagao T., Molchanov O., Hattori K., Orihara Y., Gotoh K., Akinaga Y., Tanaka H. (2002). Electric and magnetic phenomena observed before the volcano-seismic activity in 2000 in the Izu Island Region, Japan. *Proceedings of the National Academy of Sciences of the United States of America*, 99, 7352–7355. [This paper presents significant anomalous changes in the ultra low frequency range observed in both geoelectric and geomagnetic fields before the major volcano-seismic activity in the Izu Island region, Japan.]

Varotsos P., Alexopoulos K., Nomicos K., Lazaridou M. (1986). Earthquake predictions and electric signals. *Nature*, 322, 120. [This is brief communication regarding the applications of the VAN method for earthquake prediction.]

Ward S.N. (1992). An application of synthetic seismicity in earthquake statistics: The Middle America Trench, *Journal of Geophysical Research*, 97, 6675–6682. [This paper demonstrates how synthetic seismicity calculations, which are based on the concept of fault segmentation and incorporate the physics of faulting through static dislocation theory, can improve earthquake recurrence statistics and hone the probabilities of hazard.]

Ward S.N. (2000). San Francisco Bay Area earthquake simulation: A step toward a standard physical earthquake model, *Bulletin of the Seismological Society of America*, 90, 370–386. [This article represents a first step in developing a standard physical earthquake model for the San Francisco Bay Area through realistic, 3000-year simulations of earthquakes on all of the area's major faults.]

Wells D.L., Coppersmith K.J. (1994). New empirical relationships among magnitude, rupture length, rupture width, rupture area, and surface displacement, *Bulletin of the Seismological Society of America*, 84, 974–1002. [Source parameters for historical earthquakes worldwide are compiled to develop a series of empirical relationships among moment magnitude, surface rupture length, subsurface rupture length, down-dip rupture width, rupture area, and maximum and average displacement per event.]

Widiyantoro S., van der Hilst R. (1996). Structure and evolution of lithospheric slab beneath the Sunda arc, Indonesia. *Science*, 271, 1566–1570. [This paper presents seismic tomography results revealing anomalies beneath the Sunda island arc, Indonesia, that suggest that the lithospheric slab penetrates to a depth of at least 1500 kilometers.]

Working Group on California Earthquake Probabilities (1988). Probabilities of large earthquakes occurring in California on the San Andreas fault, U.S. Geological Survey Open-File Report 1988-398. [This report presents the assessment of the probabilities for large earthquakes resulting from slip on the major faults of the San Andreas fault system. The evaluations are based on a probability model that assumes increase of probability with elapsed time since the previous major earthquake on the fault segment.]

Working Group on California Earthquake Probabilities (2007). The Uniform California Earthquake Rupture Forecast, Ver. 2 (UCERF 2). U.S. Geological Survey Open-File Report 2007-1437. [This report describes a new earthquake rupture forecast for California developed by the 2007 Working Group on California Earthquake Probabilities (WGCEP 2007).]

Wyss M. (ed.) (1991). *Evaluation of Proposed Earthquake Precursors*, 94 pp. Washington, D.C., American Geophysical Union. [This volume discusses earthquake precursors and their ability to forecast a target earthquake.]

Yikilmaz M.B., Turcotte D.L., Yakovlev G., Rundle J. B., Kellogg L. H. (2010). Virtual California earthquake simulations: simple models and their application to an observed sequence of earthquakes. *Geophysical Journal International*, 180, 734–742. [In this paper, the Virtual California earthquake simulation model is applied to simple problems involving a straight strike-slip fault.]

Zaliapin I., Keilis-Borok V., Ghil M. (2003). A Boolean delay model of colliding cascades. II: Prediction of critical transitions. *Journal of Statistical Physics*, 111, 839–861. [Prediction of abrupt overall changes in the behavior of hierarchical complex systems is considered in the paper on the basis of the model, which merges the physical concept of colliding cascades with the mathematical framework of Boolean delay equations.]

Zechar J.D., Jordan T.H. (2008). Testing alarm-based earthquake predictions. *Geophysical Journal of International*, 172, 715–724. [The paper presents a method for testing alarm-based earthquake predictions, which is based on the Molchan diagram (a plot of miss rate and fraction of space-time occupied by alarm) and is applicable to a wide class of predictions.]

Zhang-li C., Pu-xiong L., De-yu H., Da-lin Z., Feng X., Zhi-dong W. (1984). Characteristics of regional seismicity before major earthquakes. In: *Earthquake Prediction*, Paris, UNESCO, pp. 505-521. [This article discusses the characteristics of seismicity prior large earthquake in China.]

Zöller G., Hainzl S. (2007). Recurrence time distributions of large earthquakes in a stochastic model for coupled fault systems: the role of fault interaction. *Bulletin of the Seismological Society of America*, 97, 1679–1697. [This paper presents the study of the effect of fault interaction on the recurrence time distribution of large earthquakes on the same fault.]

Zöller G., Hainzl S., Holschneider M., Ben-Zion Y. (2005). Aftershocks resulting from creeping sections in a heterogeneous fault. *Geophysical Research Letters*, 32, L03308, doi:10.1029/2004GL021871. [This paper presents a model consisting of brittle fault segments separated by creeping zones generates realistic aftershock sequences with space-time characteristics compatible with observations.]

### **Biographical Sketch**

**Alik Ismail-Zadeh** is a Senior Scientist at the Karlsruhe Institute of Technology (KIT), Germany; Chief Scientific / Research Professor at the Russian Academy of Sciences in Moscow; and Visiting Professor at the Institut de Physique du Globe in Paris, France. He graduated from the Baku State University, Azerbaijan (mathematics, 1982) and the Lomonosov Moscow State University in Russia (mathematics and physics, 1983) before being awarded the Ph.D. (1990) and the D.Sc. (1997), both in geophysics, from the Russian Academy of Sciences. He lectures at KIT and Abdus Salam International Centre for Theoretical Physics (Trieste, Italy) and worked at several universities as visiting scholar/professor, including University of California, Los Angeles (USA); University of Cambridge (UK); University of Tokyo (Japan); University of Trieste (Italy); University of Uppsala and Royal Institute of Technology (Sweden). Scientific interests of Alik Ismail-Zadeh cover studies of dynamics of the crust and upper mantle and their surface manifestations (including seismicity, seismic hazard and risk, sedimentary basins, salt tectonics) through multidisciplinary synthesis, theoretical and computational analysis. He is a principal author and coauthor of over 90 peer-reviewed papers and four books. Alik Ismail-Zadeh is Secretary-General of the International Union of Geodesy and Geophysics (IUGG), Chair of the Natural Hazards Focus Group of the American Geophysical Union (AGU), and Past-President of the IUGG GeoRisk Commission. He is a recipient of the Academia Europaea Young Scientist Award and the AGU International Award, and he was awarded several research fellowships including the Russian President, the Alexander von Humboldt Foundation, the Royal Society of London, and the Royal Swedish Academy of Sciences.