

GEOMORPHOLOGY AND BIOGEOGRAPHY OF TROPICAL DESERTS

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Summary

Geomorphology of Tropical Deserts is driven by the dry climatic condition of the environment. The heat provided by the sun reaches the surface and creates a high temperature in days, but at night the temperature falls. This cycle creates the conditions conducive to the physical weathering that predominates in the tropical deserts. Aeolian processes commands the shaping of surfaces, moving sediments provided by the weathering. These conditions occur in a specific area of the Earth near the sub-tropical

areas. The fauna and vegetation of Tropical Deserts are adapted to the extreme environmental conditions. Adaptive processes were developed to suit their supplies of water and food.

1. Introduction

Deserts ecosystems located in tropical regions are environment of extremes, with lack of moisture and generally synonymous with arid regions. They are one of the hottest and driest areas of the planet, with no or sporadic rainfall. These conditions are due to some different conditions as continentality, topography and subtropical high pressure cells influence. These kinds of deserts are located in south and north hemispheres, especially between 5 and 30 degrees of latitude.

The tropical desert is an environment of extremes: it is the driest and hottest place on earth. Rainfall is sporadic and in some years no measurable precipitation falls at all. The terribly dry conditions of the deserts are due to the year-round influence of subtropical high pressure and continentality.

In tropical areas the heat enhances evaporation and the dryness conditions of the areas with little precipitation. Rain also occurs in a few events and quickly the moisture is absorbed by the soil or evaporated. These climatic conditions do not allow the geochemical processes of weathering to happen and most of the rock transformations are due to physical processes of contraction and expansion with the break of rocks in fragments. In this case, most of the desert surface is occupied by fragmented rocks, outcrop rocks or sand. Lack of vegetation, bare rock or sand, mountains and canyons, mesas or buttes, dunes, basins and playas, washes and arroyos, alluvial fans and generally angular topography are all deserts characteristics.

Geomorphic processes performed by the wind are called eolian processes, or aeolian after the Greek god of the wind Aeolus. In tropical deserts the work of wind shapes the Earth's surface. Though dominating dry climates, eolian processes are also effective in semi-arid, sub-humid or humid regions as well.

2. Tropical Deserts

2.1. Geographical Distribution

Tropical Deserts are typically found in continental interiors of the tropics with extensions to the subtropical areas and on the leeward side of mountains. Such deserts are also found in cool coastal regions where cold water upwells along a coast, stabilizing the air and preventing moisture formation like that near Chile's coast. Vast deserts cover parts of Africa – The Kalahari, The Sahara, extensive areas in the Middle East from Saudi Arabia to Iran, part of south Asia in Pakistan and western India. Tropical deserts are found also in Baja California and interior Mexico in North America. In South America this kind of deserts occurs in north of Chile and south parts of Bolivia and Peru, The Atacama. (Figure 1 – Table 1)

Desert's Name	Area (km ²)	Country/Countries
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Sahara Desert	8 600 000	Egypt, Libya, Chad, Mauritania, Morocco, Algeria.
Arabian Desert	2 331 000	Saudi Arabia, Jordan, Iraq, Iran, Kuwait, Qatar, United Arab Emirates, Oman and Yemen.
Great Victoria Desert	647 000	Australia
Chihuahuan Desert	450 000	Mexico and United States
Great Sandy Desert	400 000	Australia
Kalahari Desert	260 000	Botswana, Namibia and South Africa
Syrian Desert	260 000	Syria, Jordan and Iraq
Gibson Desert	155 000	Australia
Simpson Desert	145 000	Australia
Atacama Desert	140 000	Chile and Peru
Namib Desert	135 000	Namibia

Table 1. Major tropical desert by area.

Source: http://en.wikipedia.org/wiki/List_of_deserts_by_area.

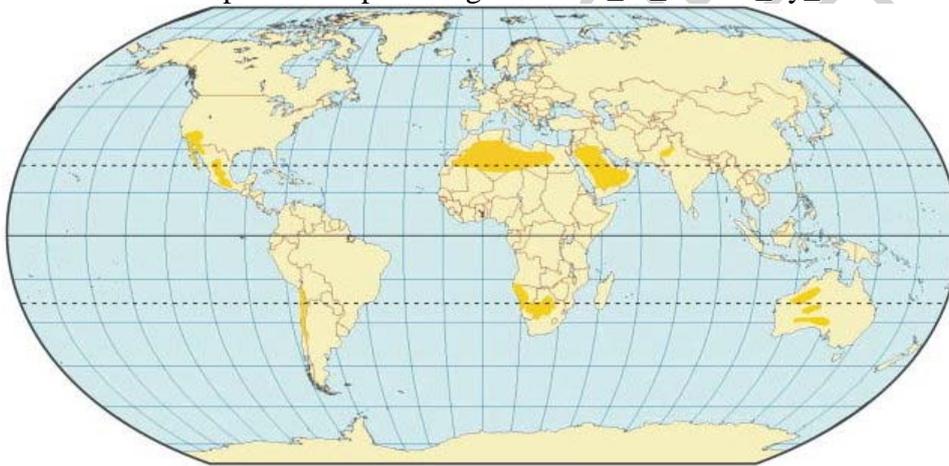


Figure 1. Location of tropical deserts.

2.2. Controlling Factors

The most important controlling factor for the tropical desert climate is the year-round presence of subtropical high pressure with hot, dry descending air cells called Hadley Cells, the Sahara, Arabian, Sonora and northern Atacama Desert are of this type. The effect of descending air from subtropical high adiabatically warms causing the air to dry out and inhibit condensation. Also, if the distance from moisture sources increases, aridity will also rise. Situation in the mountainous areas which provokes rain shadow conditions also promotes dry conditions. Cool coastal deserts are found in areas along coasts where cold water is upwelling them.

Deserts occur in specific latitudes (5-35° north and south of the equator) because of the general thermodynamics of our planet. Solar radiation hits the earth with highest intensity near the equator. Because the earth's axis is tilted 3.5° with respect to the plane of its orbit, during part of the year the zone of maximum solar interception shifts

northwards, towards the Tropic of Cancer, and during part of the year it moves southwards, towards the Tropic of Capricorn. Thus, the warm tropics form a belt around the equator from latitude 3° north to latitude 3° south called Intertropical Convergence Zone, where the tropical heat generates rising, unstable air.

When the air moves up, it condenses the moisture evaporated from seas and forests at warm tropical areas. In this movement from the equator to the tropics at high altitudes, the air cools again and starts descending towards the midlatitudes both north and south. The air masses become extremely dry in this movement because of the loss of their moisture during their tropical ascent. So, the mid-latitude arid fringes that run alongside the tropical belts have a more stable atmosphere than the equatorial areas.

2.3. Distinguishing Characteristics

2.3.1. Temperature

Tropical deserts present the highest average annual temperature of any climate on the planet. These high temperatures are motivated by the highest percentage of sunshine of any climate caused by the high sun angles throughout the year. The lowest temperature never reaches below 18°C and many places have consecutive average monthly temperatures in the mid 30 degrees Celsius. The high energy input during the day and large loss at night results in an extremely large daily temperature range. During days the temperatures at low elevation inland deserts can reach 40-50°C and the daily difference could reach 50°C.

The tropical desert sky remains cloud-free due to the subsiding air of dominant high pressure resulting in large amounts of insolation. The high insolation days caused by the cloudless sky not only causes an intensive amount of energy on the land surface, but also lets much heat out at night.

2.3.2. Precipitation

The precipitation is very irregular in the tropical deserts. Deserts located at low latitude have an average precipitation less than 250 mm.y⁻¹. Some desert have years without rains, because they are located far from the ocean and in areas over intense action of high pressure cells, a condition that combines and creates a very dry environmental condition.

The tropical desert climate is influenced by upper air stability and subsidence which is the result of the presence of the subtropical high pressure zone. Relative humidity is normally low, averaging 10 to 30 percent in interior locations. Precipitation is very low in quantity and very infrequent in distribution, both temporally and spatially. Temperature varies greatly both diurnally and annually. The highest average monthly temperatures on the Earth are found in the tropical desert. They range between 29 to 35 degrees Celsius. Winter monthly temperatures can be 15 to 25 degrees cooler than summer temperatures. This climate also has extreme diurnal ranges of temperature. The average diurnal range is from 14 to 25 degrees Celsius. (Table 2).

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Biographical Sketches

Sílvio Carlos Rodrigues graduated in 1987 from the Universidade de São Paulo and obtained his doctoral degree from the same University in 1998. He had worked for some consulting enterprises from 1988 to 1999, especially on Environmental Impact Assessments of hydroelectric power stations in large rivers of the Amazonian and Savanna regions, and more recently on the Economic-Ecologic Management of Mato Grosso and Rondonia states. After 1999 he joined the Universidade Federal de Uberlândia (UFU), where he is currently employed. Sílvio Rodrigues is coordinator of the Soil Erosion and Geomorphology Laboratory at UFU, where he and his students carry out research on soil erosion and geomorphological cartography. Current projects include measurements of gully evolution and experiments on sheet erosion. From August 2004 until December 2006 he was President of Brazilian Geomorphological Union, which organizes the Brazilian community of geomorphologists and related researchers, including soil scientists and quaternary geologists. Since 2003 he has been member of the COMLAND – Commission on Land Degradation and desertification of the International Association of Geography.

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